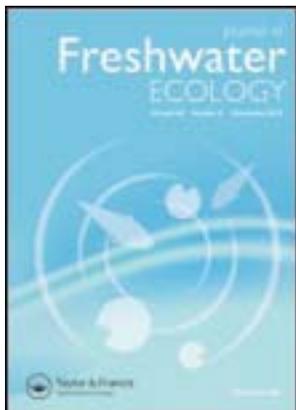


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## Comparison of $^{15}\text{N}_2$ flux and acetylene inhibition denitrification methods in Atchafalaya River basin sediments

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Flooded swamps in the Atchafalaya River basin, Louisiana, USA, are thought to be a major mechanism for removing excess riverine nitrate that may be causing hypoxia in coastal Louisiana. Two denitrification methods, the  $^{15}\text{N}_2$  flux and acetylene inhibition technique, were compared and evaluated. Sediment samples were collected from three baldcypress swamp sites located within the Atchafalaya River basin. Highly  $^{15}\text{N}$ -labeled (61.7 at.%) and unlabeled nitrate ( $100 \text{ mg NO}_3\text{-N L}^{-1}$ ) were applied to the floodwater of separate sets of sediment cores. Denitrification potential was measured 2, 6, 10, and 16 days after nitrate addition. Two days after nitrate addition, no significant differences in denitrification rates were observed between methods. On days 6, 10, and 16, denitrification measured by the  $^{15}\text{N}_2$  method was higher compared to the acetylene inhibition procedure. When significance was observed,  $^{15}\text{N}_2$  denitrification was approximately 5.3 times greater.

**Keywords:** denitrification;  $^{15}\text{N}_2$  flux method; acetylene inhibition technique; Atchafalaya River Basin

### Introduction

The Mississippi–Atchafalaya River basin (MARB) drains approximately 41% of the conterminous United States (all or part of 31 states). Nitrogen fertilizer applied to MARB farmland has increased from less than one million metric tons in 1951 to about seven million metric tons in 1996 (Goolsby et al. 2001).

Approximately 90% of the annual discharge of nitrogen (N) to the Gulf of Mexico (GOM) is transported by the Mississippi River and Atchafalaya River (Dunn 1996). Average annual flux of N (1981–2005) to the GOM was 1,470,000 metric tons of which 61% was inorganic nitrate (Aulenbach et al. 2007). About 90% of the nitrate entering the GOM emanates from nonpoint pollution sources (mainly agriculture), with highest discharges observed in the spring and early summer.

Excess N (mainly nitrate) from the MARB may be one of the major causes of low dissolved oxygen ( $<2 \text{ mg L}^{-1}$ ) in a large hypoxic area along the Louisiana–Texas coasts (Rabalais et al. 2002). Hypoxia can cause increased stress or death to coastal

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bottom-dwelling organisms that cannot flee the oxygen depleted zone. The size of the GOM hypoxic zone is highly variable and has increased over the past two decades. The zone averaged approximately 8300 km<sup>2</sup> from 1985 to 1992, but from 1993 to 2001 averaged 16,000 km<sup>2</sup> (Rabalais et al. 2002). The size of the hypoxic zone in 2007 was estimated at 20,500 km<sup>2</sup> and may increase over the next few years due to increased ethanol production from corn in the MARB (USGS 2007).

The Mississippi River/Gulf of Mexico Watershed Nutrient Task Force (2001) developed an action plan to reduce hypoxia and the size of the zone to less than 5000 km<sup>2</sup> by 2015 for the northern GOM. Enhancing and/or restoring natural denitrification has been proposed to reduce nitrogen loads. Remediation actions to increase denitrification include restoring and creating riparian buffers and coastal wetlands and diverting nitrate-elevated river water into Louisiana's coastal ecosystems (USEPA 2001; Mississippi River/Gulf of Mexico Watershed Nutrient Task Force 2008).

Denitrification is the microbial reduction of nitrate (NO<sub>3</sub><sup>-</sup>) and nitrite (NO<sub>2</sub><sup>-</sup>) to gaseous end products nitric oxide (NO), nitrous oxide (N<sub>2</sub>O), and dinitrogen (N<sub>2</sub>) which represents a major pathway for N removal from aquatic and wetland ecosystems (Mitsch and Gosselink 2000). Denitrification processes are affected by soil moisture and texture, pH, microbial community structure, sediment O<sub>2</sub>, NO<sub>3</sub><sup>-</sup> concentration, organic matter, and temperature (Groffman 1994; Clement et al. 2002; Seitzinger et al. 2006; Hernandez and Mitsch 2007). Denitrification is fueled by NO<sub>3</sub><sup>-</sup> produced from mineralization of inorganic NH<sub>4</sub><sup>+</sup> and by water column NO<sub>3</sub><sup>-</sup> diffusion into reduced sediments.

Direct (measurement of N<sub>2</sub>O and/or N<sub>2</sub>) and indirect procedures have been used to estimate denitrification. Denitrification methods include N<sub>2</sub> quantification, <sup>15</sup>N stable isotope tracers, N<sub>2</sub>:Ar ratios, acetylene inhibition, nitrogen mass balance, nitrate disappearance, molecular techniques, and stoichiometric approaches. All denitrification methods have advantages, disadvantages, limitations, and assumptions that must be evaluated prior to method selection (Cornwell et al. 1999; Steingruber et al. 2001; Groffman et al. 2006; Wallenstein et al. 2006).

The objective of this investigation was to compare and evaluate two denitrification methods after NO<sub>3</sub><sup>-</sup> addition to flooded sediments. Potential denitrification rates were measured using the <sup>15</sup>NO<sub>3</sub><sup>-</sup> → <sup>15</sup>N<sub>2</sub> flux method and the acetylene inhibition technique.

### Materials and methods

Baldcypress sediments were collected from three flooded locations within the Atchafalaya River basin. Sites were located along Bayou Cowan (BC), Jones Bayou (JB), and Bee Bayou (BB). Bulk sediment (0–15 cm depth) was manually collected, sealed in polyethylene bottles, and iced. Sediments were transported to the laboratory and stored at 1–2 C. Total sediment organic matter was measured by loss of ignition at 550 C for 2 h and total N was determined with a Leco C and N analyzer (Leco Corp., St. Joseph, Michigan, USA). Clay content was determined by the pipette method. Sediment organic matter, total N, and clay content ranged from 16.8 to 32.3, 0.4 to 0.8, and 53% to 66%, respectively. The lowest concentrations of these components were observed for BB, and highest were for BC. Substrate pH values varied from 4.4 (JB) to 5.5 (BB).

For each sediment type (site), six glass incubation jars (9 cm diameter by 16 cm height) were filled with sediment to a height of 6 cm, and the sediment was then packed to field bulk density. A 4-cm layer of nitrate-free distilled water was added, and cores were allowed to incubate (22 C) for 1 week to re-establish the thin surface oxidized layer (observed in the field) overlaying the reduced zone. Then, 180 mg of  $^{15}\text{N}$  labeled (61.7 at.%)  $\text{KNO}_3$  (dissolved) was added to the floodwater of two sediment cores, and 180 mg of dissolved  $\text{KNO}_3$  (unlabeled) was added to the floodwater of the remaining four cores. A high concentration of  $\text{NO}_3\text{-N}$  was required to assure sediment cores were not nitrate-limited and to detect  $^{15}\text{N}_2$  in the headspace volume (containing 78%  $\text{N}_2$ ) at the end of the 6-hour incubation. Labeled nitrate was only added to two cores due to cost. Gas samples were collected at 2, 6, 10, and 16 days after nitrate addition.

#### *Acetylene inhibition method*

The acetylene ( $\text{C}_2\text{H}_2$ ) inhibition method was used to estimate potential denitrification (Groffman 1994). After unlabeled nitrate (dissolved) was added to the floodwater, the incubation jars were sealed (screw caps fitted with rubber septa), and purified  $\text{C}_2\text{H}_2$  was then added to the headspace volume (10% v/v) and floodwater. Headspace gas samples were collected 2 and 6 h after  $\text{C}_2\text{H}_2$  addition with a 2-mL gas-tight syringe to estimate the increase of  $\text{N}_2\text{O}$ . Two hours were allowed for  $\text{C}_2\text{H}_2$  diffusion through the floodwater into the sediment before gas collection. After 6 h, screw caps were removed and cores were exposed to the atmosphere. Floodwater/sediment cores were only exposed to  $\text{C}_2\text{H}_2$  for a short time to minimize effects on microorganisms (Watts and Seitzinger 2000). Nitrous oxide fluxes were calculated using the closed chamber equation of Rolston (1986):  $F = (V/A)(273/T)(\Delta C/\Delta T)$ , where V is core headspace volume, A the core sediment area, T the absolute temperature of headspace gases, and  $\Delta C/\Delta T$  the change in  $\text{N}_2\text{O}$  concentration per unit of time. The Bunsen absorption coefficient was used to estimate  $\text{N}_2\text{O}$  dissolved in the floodwater. Nitrous oxide flux is reported as  $\text{mg N evolved m}^{-2} \text{d}^{-1}$  which represents total denitrification ( $\text{N}_2\text{O} + \text{N}_2$ ).

Nitrous oxide concentrations were determined with a Shimadzu GC-14A gas chromatograph (Shimadzu Scientific Instruments, Inc., Columbia, Maryland, USA) fitted with a 1-mL sample loop, Poropak Q 1.8 m ss column, and electron capture detector (ECD) and calibrated with certified  $\text{N}_2\text{O}$  reference standards (Scott Specialty Gases, Inc., Plumsteadville, Pennsylvania, USA). Nitrogen was the carrier gas and the instrument operated at 40, 100, and 290 C for the oven, injector, and ECD, respectively (Lindau et al. 1998).

#### *$^{15}\text{N}_2$ method*

Highly  $^{15}\text{N}$ -labeled nitrate was added to the floodwater of two sediment cores collected from each of the three sites. Incubation jars were sealed and headspace gas samples were collected at 0 (background) and 6 h with a gas-tight syringe. Collected gas samples were immediately transferred into evacuated glass Vacutainers (100 mm length by 16 mm i.d.). Prior to gas collection, Vacutainers (Becton Dickinson, Franklin Lakes, New Jersey, USA) were placed on a high-vacuum system to remove residual gases. A slight over-pressure of headspace gas was injected into each

Vacutainer to prevent atmospheric gas contamination. Screw caps were removed at the end of each 6 h incubation.

Stable  $^{15}\text{N}_2$  isotopic distribution (28, 29, and 30 masses) was determined with a gas chromatograph coupled to a triple collector 20/20 isotope ratio mass spectrometer (GC-IRMS, Robo Prep-G Plus, Europa Scientific, SerCon Ltd., UK). Fluxes of  $^{15}\text{N}_2$  from the floodwater/sediment cores were calculated using the equations and  $^{15}\text{X}_\text{N}$  and  $d$  values outlined by Hauck and Bouldin (1961) and Mulvaney and Boast (1986). Emission of  $^{15}\text{N}_2$  is reported as  $\text{mg N evolved m}^{-2}\text{d}^{-1}$ .

### Statistical methods

All statistical analyses were conducted using SAS 9.1.3 (SAS Institute Inc., Cary, North Carolina). Denitrification data was evaluated by the analysis of variance (ANOVA) test procedure and significant differences between treatment means were calculated by Duncan's multiple range test. An alpha level of 0.05 was used for all statistical analyses.

### Results

Denitrification rates measured by the  $^{15}\text{N}_2$  method generally increased (after  $\text{NO}_3$  addition) from day 2 to day 10 and then decreased on day 16. Denitrification potential measured by the  $\text{C}_2\text{H}_2$  technique steadily decreased over the 16 day incubation. On days 6, 10, and 16,  $^{15}\text{N}_2$  denitrification fluxes were higher compared to  $\text{C}_2\text{H}_2$  treatment means (Figure 1).

BC  $^{15}\text{N}_2$  denitrification ranged from 17.1 to 26.7  $\text{mg N evolved m}^{-2}\text{d}^{-1}$  and averaged 21.3  $\text{mg N m}^{-2}\text{d}^{-1}$  over the 16-day incubation. With the  $\text{C}_2\text{H}_2$  procedure, BC denitrification ranged from 0.8 to 35.7  $\text{mg N m}^{-2}\text{d}^{-1}$  with an average flux of 12.6  $\text{mg N m}^{-2}\text{d}^{-1}$ . Over the entire incubation,  $^{15}\text{N}_2$  denitrification was approximately 1.7 times greater. On day 2, denitrification potential  $^{15}\text{N}_2$  and  $\text{C}_2\text{H}_2$  flux means were not significantly different. On days 6, 10, and 16, denitrification measured using the  $^{15}\text{N}_2$  method was significantly higher (Figure 1).

Denitrification measured on BB sediment varied from 15.4 to 39.1  $\text{mg N m}^{-2}\text{d}^{-1}$  and averaged 28.0  $\text{mg N m}^{-2}\text{d}^{-1}$  ( $^{15}\text{N}_2$  method). Potential denitrification ranged from 4.4 to 14.4  $\text{mg N m}^{-2}\text{d}^{-1}$  with an average of 9.3  $\text{mg N evolved m}^{-2}\text{d}^{-1}$  using the  $\text{C}_2\text{H}_2$  method (Figure 1). Dinitrogen fluxes from  $^{15}\text{N}$  cores were about 3.0 times higher compared to  $\text{C}_2\text{H}_2$  method denitrification means. On days 6, 10, and 16,  $^{15}\text{N}_2$  denitrification means were significantly higher, but on day 2, no significant differences were observed between the methods.

Using the  $^{15}\text{N}_2$  method, JB denitrification rates varied from a low of 16.5 to a maximum of 46.7  $\text{mg N m}^{-2}\text{d}^{-1}$  and averaged 29.9  $\text{mg N m}^{-2}\text{d}^{-1}$  over the sampling period. Acetylene inhibition denitrification means were lower and ranged from 3.3 to 20.3  $\text{mg N m}^{-2}\text{d}^{-1}$  (Figure 1) and averaged 11.2  $\text{mg N m}^{-2}\text{d}^{-1}$  over the 16 days.  $^{15}\text{N}$ -labeled denitrification values were approximately 2.7 times greater compared to the  $\text{C}_2\text{H}_2$  procedure. JB denitrification fluxes measured by the  $^{15}\text{N}_2$  method were significantly higher on days 10 and 16.

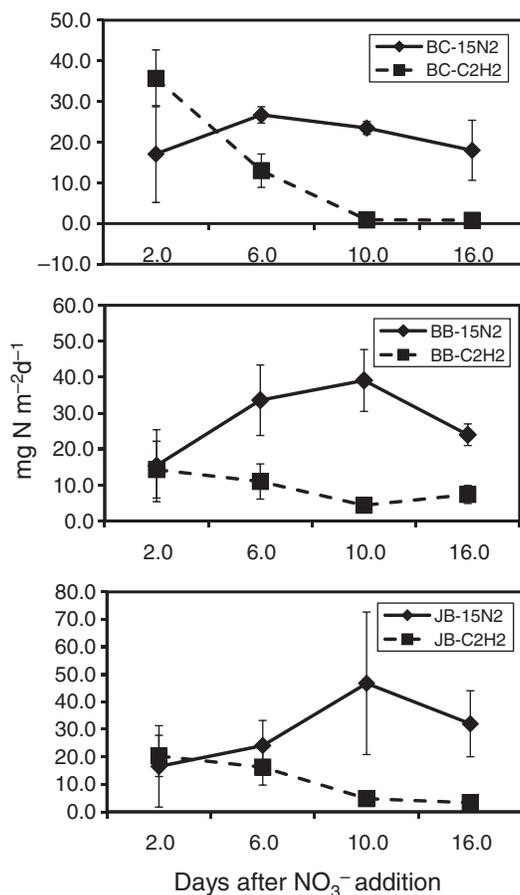


Figure 1. <sup>15</sup>N<sub>2</sub> and C<sub>2</sub>H<sub>2</sub> methods potential denitrification means (mg N m<sup>-2</sup> d<sup>-1</sup>) for Bayou Cowan (BC), Bee Bayou (BB), and Jones Bayou (JB) days after NO<sub>3</sub><sup>-</sup> addition to sediment cores. Error bars are one standard deviation.

### Discussion

Across locations, denitrification rates measured by the <sup>15</sup>N<sub>2</sub> flux and C<sub>2</sub>H<sub>2</sub> methods were not significantly different 2 days after nitrate addition. For the remaining days (6, 10, and 16), <sup>15</sup>N<sub>2</sub> denitrification was significantly higher eight out of nine sampling times and was approximately 5.3 times greater than C<sub>2</sub>H<sub>2</sub> method denitrification. Denitrification <sup>15</sup>N<sub>2</sub> fluxes evolved into the floodwater/sediment core headspace may significantly underestimate sediment <sup>15</sup>N<sub>2</sub> production. In a 1988 study, Lindau et al. (1988) observed approximately 34% of the applied <sup>15</sup>N was entrapped as <sup>15</sup>N<sub>2</sub> in flooded rice soil 33 days after <sup>15</sup>N addition. The <sup>15</sup>N<sub>2</sub> flux rates are a function of production rate, solubility, mass flow, and diffusion (Well and Myrold 1999).

The C<sub>2</sub>H<sub>2</sub> inhibition and <sup>15</sup>N<sub>2</sub> flux methods have limitations and assumptions that need to be considered for flooded sediments. Limitations of the C<sub>2</sub>H<sub>2</sub> technique include: inhibition of nitrification by C<sub>2</sub>H<sub>2</sub>, which can lead to underestimation of

total denitrification; incomplete blockage of  $\text{N}_2\text{O}$  to  $\text{N}_2$  by  $\text{C}_2\text{H}_2$ ; presence of hydrogen sulfide; slow diffusion of  $\text{C}_2\text{H}_2$  through the floodwater and sediment; sediment respiration enhancement; metabolism and/or decomposition of  $\text{C}_2\text{H}_2$ ; and contaminated  $\text{C}_2\text{H}_2$  (Seitzinger et al. 1993; Cornwell et al. 1999; Groffman et al. 2006). Assumptions for the  $^{15}\text{N}_2$  method include:  $^{14}\text{N}$  and  $^{15}\text{N}$  atoms are randomly distributed between  $\text{N}_2$  molecules produced during the denitrification process; total amount of  $^{28}\text{N}_2$  in core headspace does not change over the incubation period; and the  $^{15}\text{N}$  label of the added nitrate in the sediment is uniform (Siegel et al. 1982; Mulvaney and Boast 1986).

Numerous sediment denitrification methods have been compared, including:  $^{15}\text{N}$  isotope pairing and  $\text{C}_2\text{H}_2$  inhibition (Lohse et al. 1996);  $\text{N}_2$  flux and  $\text{C}_2\text{H}_2$  blockage (Watts and Seitzinger 2000);  $\text{N}_2$  flux,  $^{15}\text{N}$  isotope pairing and mass balance techniques (van Luijn et al. 1996); membrane inlet mass spectrometry and isotope pairing mass spectrometry (Smith et al. 2006); and isotope pairing and  $\text{N}_2:\text{Ar}$  ratios (Ferguson and Eyre 2007). Comparison of our denitrification methods ( $^{15}\text{N}_2$  flux and  $\text{C}_2\text{H}_2$  inhibition) to published research is difficult due to a wide variety of laboratory procedures, sediment properties, applied nitrate and ammonium concentrations, denitrification calculations, incubation methods and times, and gas collection techniques used. In a 1993 experiment, Seitzinger et al. compared three denitrification methods ( $\text{N}_2$  flux,  $^{15}\text{N}$  tracer, and  $\text{C}_2\text{H}_2$  inhibition) using aquatic sediments. Denitrification rates due to addition of nitrate to the sediment core floodwater using the  $^{15}\text{N}$  tracer method were about 35% of those measured by the  $\text{C}_2\text{H}_2$  inhibition method.

Over a short incubation time (2 days), our denitrification results measured with the  $\text{C}_2\text{H}_2$  method were comparable to the  $^{15}\text{N}_2$  flux method; however, over extended incubation periods (6, 10, and 16 days), the  $\text{C}_2\text{H}_2$  inhibition method significantly underestimated potential denitrification. Coupled nitrification–denitrification would not be captured due to  $\text{C}_2\text{H}_2$  inhibition of nitrification in sediments. In addition, slow diffusion of  $\text{C}_2\text{H}_2$  through the floodwater, incomplete blockage of  $\text{N}_2\text{O}$  by  $\text{C}_2\text{H}_2$ , and  $\text{C}_2\text{H}_2$  metabolism by microorganisms may also be responsible for underestimating denitrification. For our study, denitrification due to coupling was thought to be a minor component compared to the denitrification contribution due to the high level of  $\text{NO}_3\text{-N}$  added ( $100 \text{ mg N L}^{-1}$ ).

Further studies are needed to compare denitrification methods at Mississippi River water *in situ*  $\text{NO}_3\text{-N}$  concentrations ( $1\text{--}3 \text{ mg N L}^{-1}$ ) entering the Atchafalaya River basin and across all basin habitats. In addition, denitrification methods need to be evaluated for incubation periods less than 2 days. If the  $\text{C}_2\text{H}_2$  inhibition method can be calibrated against the  $^{15}\text{N}_2$  flux procedure, considerable expense and time can be saved on denitrification analyses. Quantification of denitrification is needed to evaluate the contribution of denitrification to nitrogen removal within the Atchafalaya River basin.

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